

**Exploring iron control on global productivity:  
“FePSi”, an inverse model of the ocean’s  
coupled phosphorus, silicon and iron cycles.**

Benoît Pasquier - Ph.D. student



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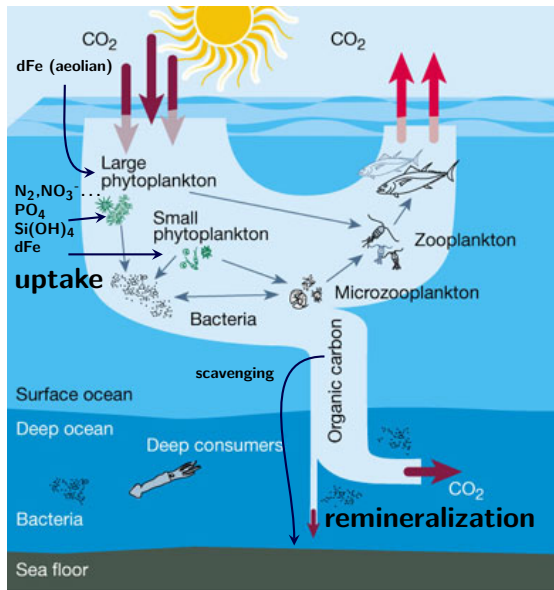
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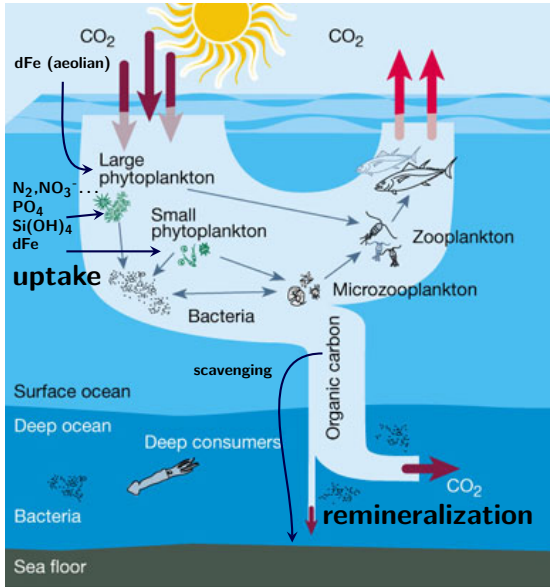
1. Can we constrain the BGC parameters of the nutrient cycles?
2. How do global nutrient cycles respond to perturbations in dFe?
3. Can we test the  $\text{Si(OH)}_4$  leakage hypothesis?

# Nutrient cycle and coupling

In the real world:



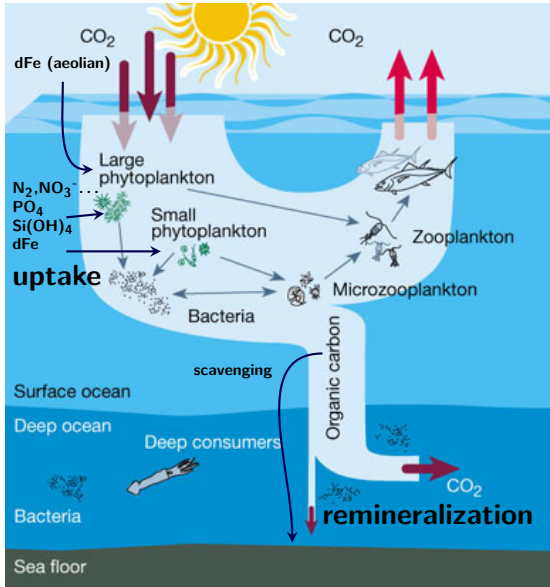
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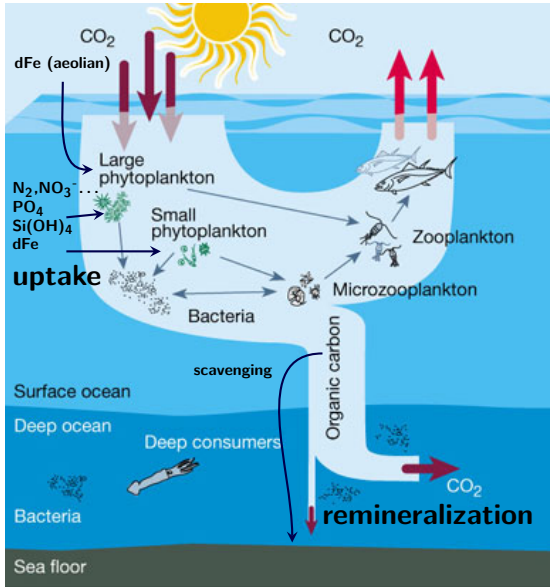


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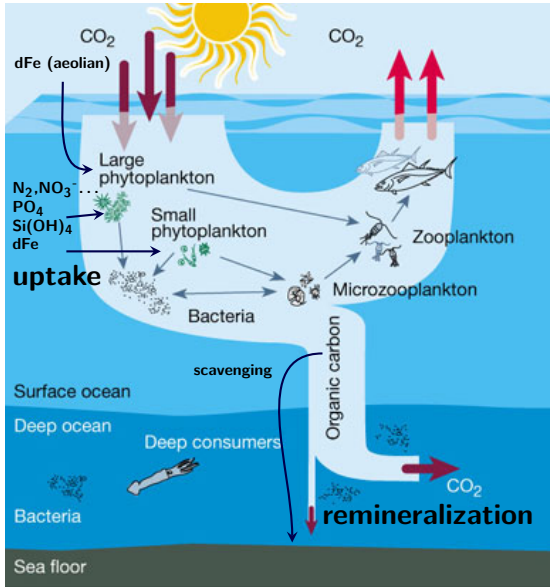
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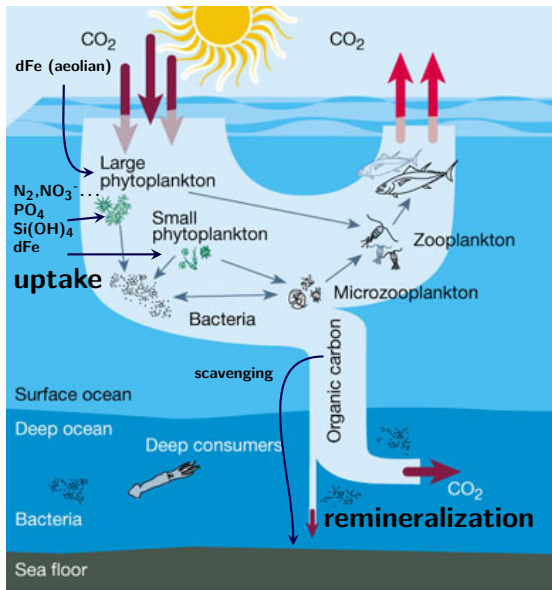
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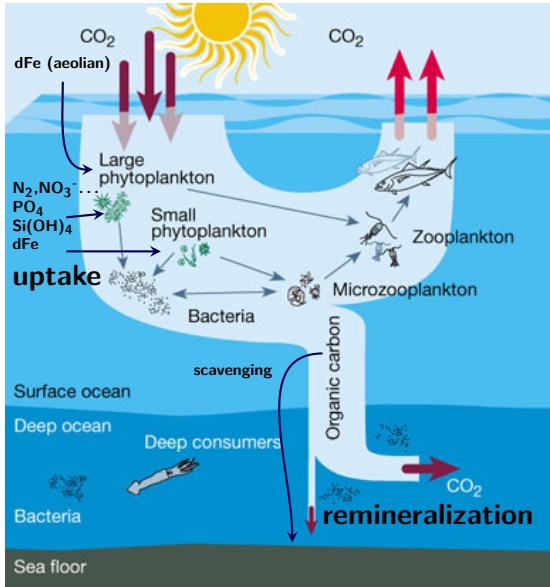


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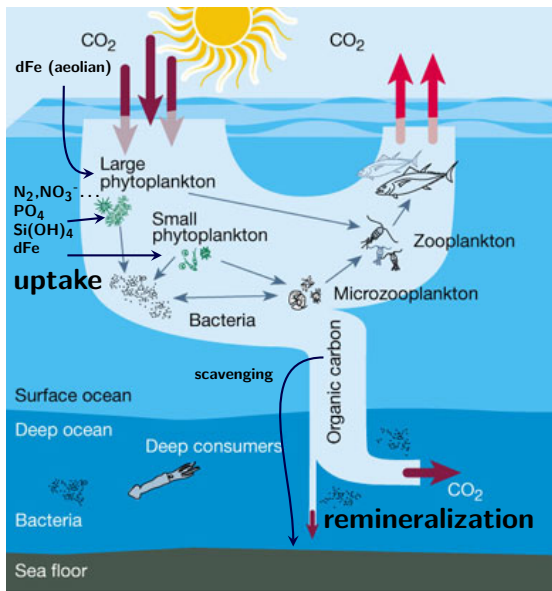
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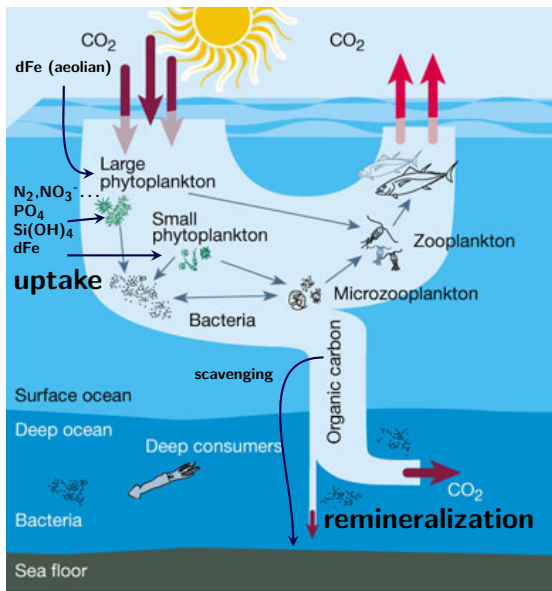
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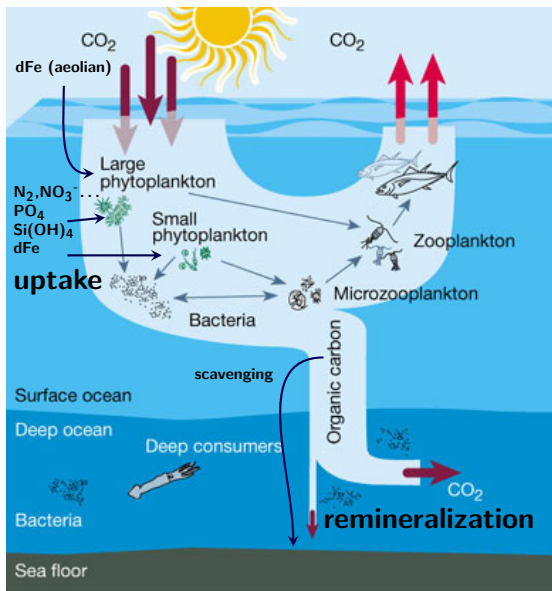
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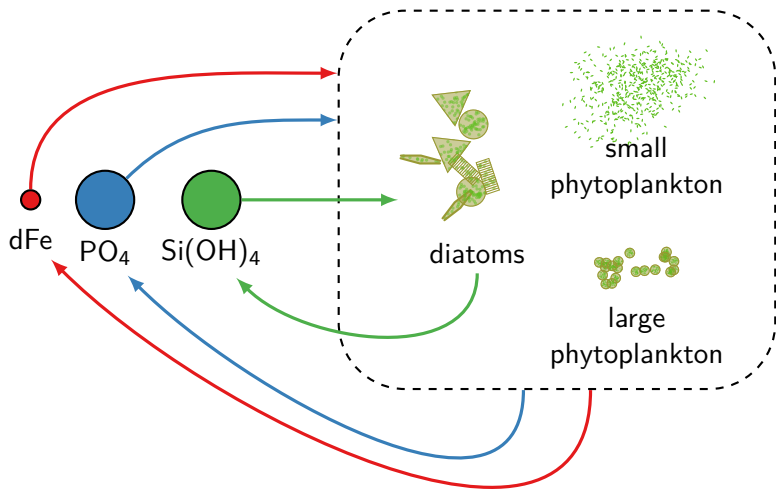
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- poor constraints on parameters
- lack of data (iron)

## Simplified biological cycling, b

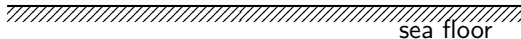


- 3 nutrients
- 3 phytoplankton classes (not transported)
- No zooplankton

*Galbraith et al., 2010;*  
*Matsumoto et al., 2013*



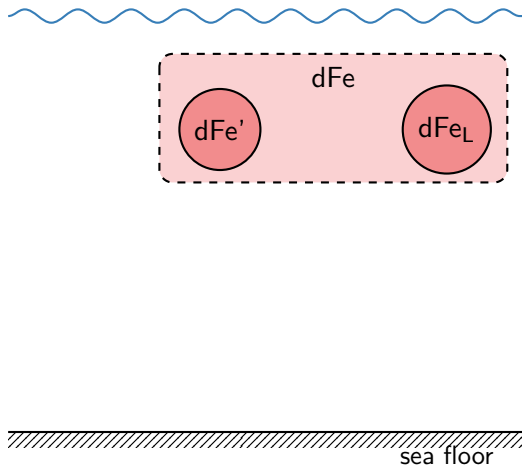
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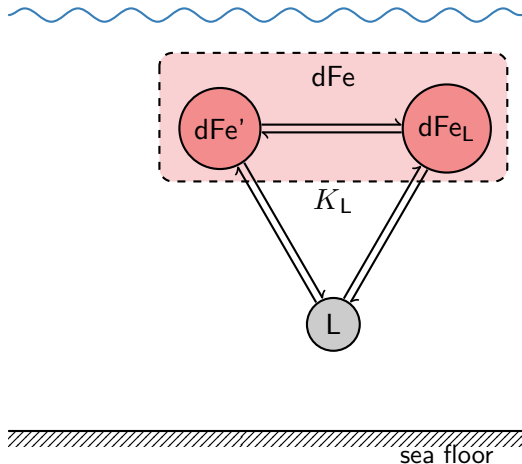
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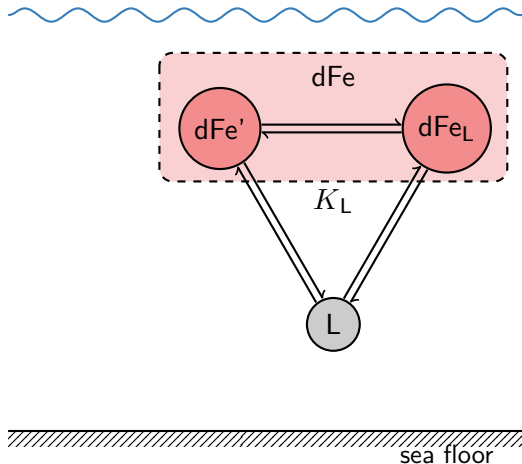
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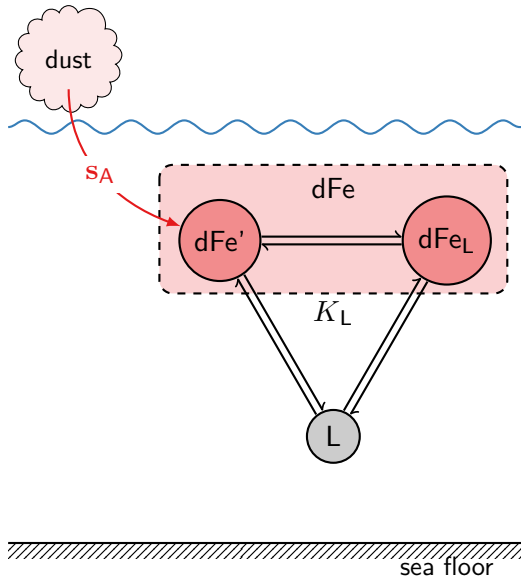
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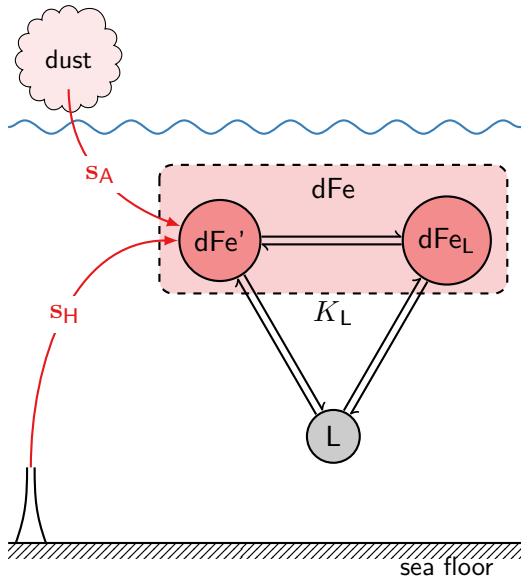
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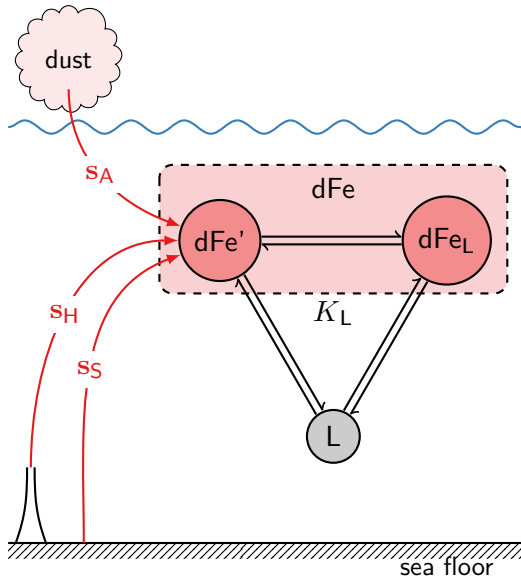
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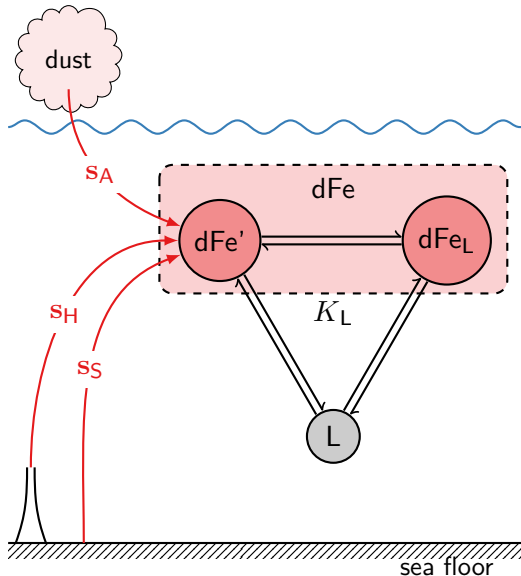
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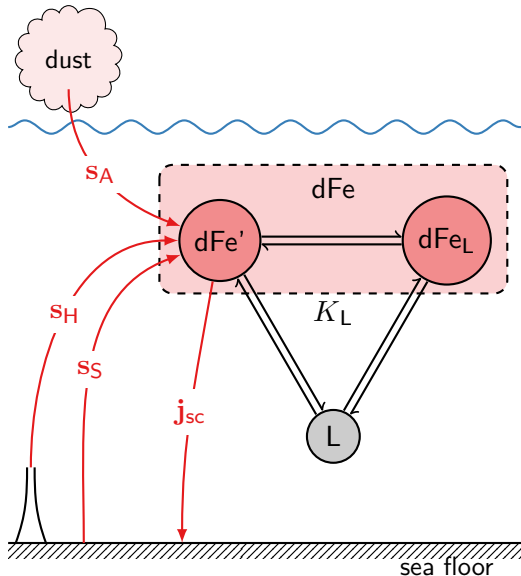


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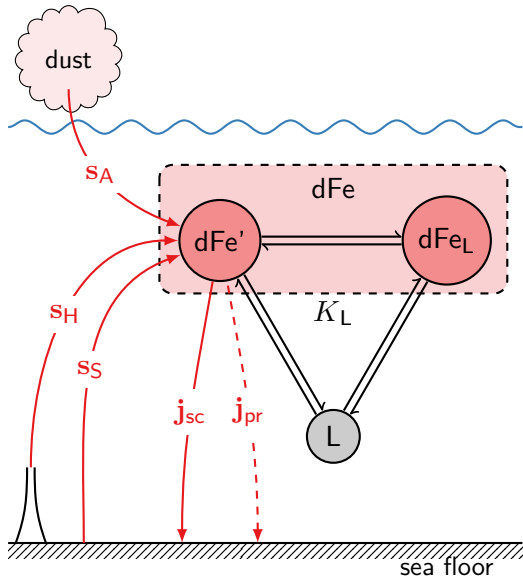
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# Discretized PDE

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- The tracer equation is  $\partial_t \mathbf{x} = \mathbf{f}(\mathbf{x})$ , where  $\mathbf{x}$  represents the (3-dimensional) concentration fields of the 3 nutrients, rearranged into a single column vector of size  $n \sim 600,000$

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- We solve the steady state equation  $\mathbf{f}(\mathbf{x}) = \mathbf{0}$  using Newton's Method



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Compare

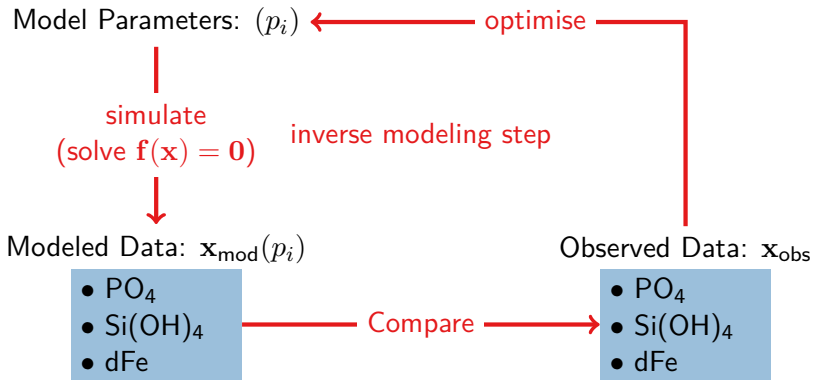
Observed Data:  $\mathbf{x}_{\text{obs}}$

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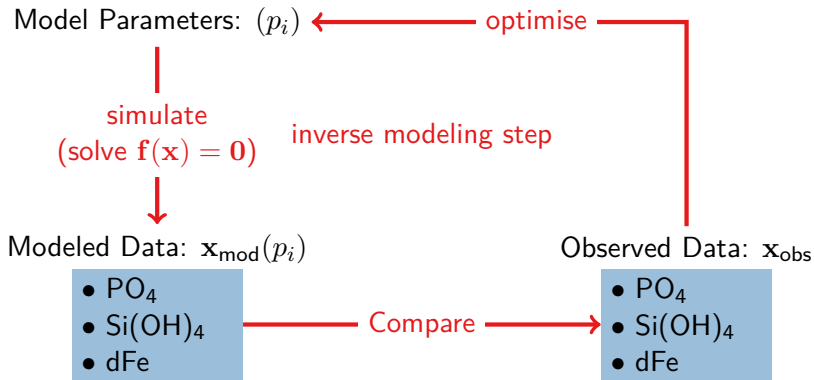
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- optimisation only possible with efficient simulation

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- Biological parameters (optimized):
  - Uptake rate timescale  $\tau$
  - Phytoplankton populations  $P_{\text{dia}}^*$ ,  $P_{\text{lrg}}^*$ , and  $P_{\text{sml}}^*$
  - (Fe:P) uptake ratio:  $R_{\text{f:p}}$  and  $k_{\text{f:p}}$
  - (Si:P) uptake ratio:  $(\text{Si} : \text{P})_{\text{min}}$ ,  $k_1$ ,  $k_2$ , and  $k_3$



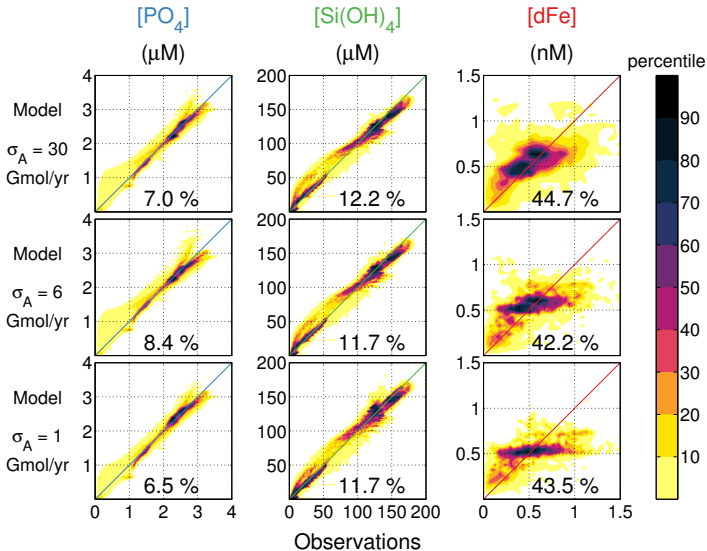
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- Iron sources and sinks (optimized for family of  $\sigma_A$ ):
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  - Inorganic scavenging by ballast particles  $f_{\text{min}}$

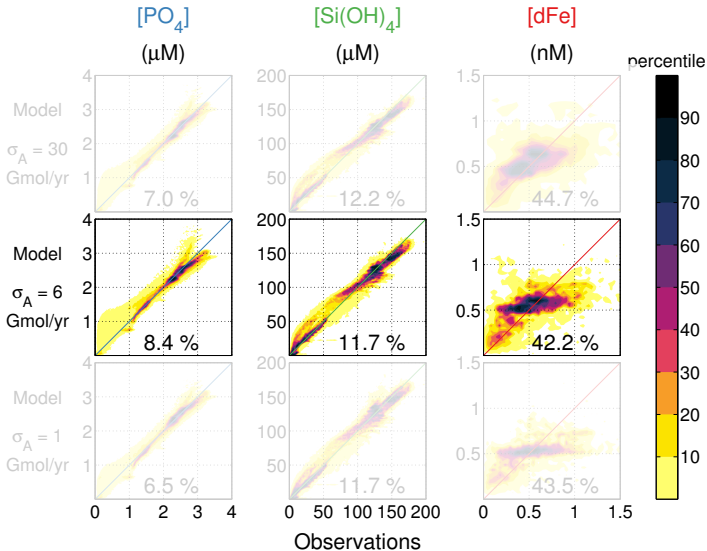
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- Non optimized (yet) parameters:
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  - local recycling fractions:  $\sigma_{\text{dia}}$ ,  $\sigma_{\text{lrg}}$ , and  $\sigma_{\text{sml}}$
  - sinking particle profiles:  $b_{\text{dia}}$ ,  $b_{\text{lrg}}$ , and  $b_{\text{sml}}$ ,
  - light harvesting efficiency:  $\alpha_{\text{min}}^{\text{chl}}$ ,  $\alpha_{\text{max}}^{\text{chl}}$ ,  $\theta_{\text{min}}^{\text{chl}}$ , and  $\theta_{\text{max}}^{\text{chl}}$
  - ligand stability constant  $K_L$

# Climatological base state - mismatch joint PDF

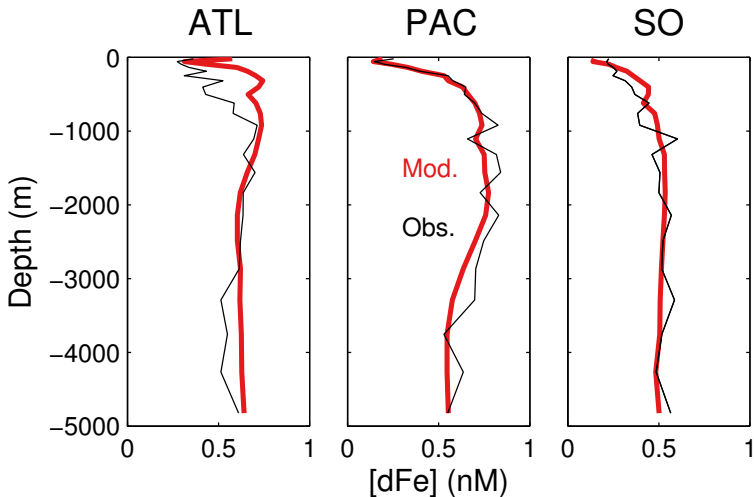


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# Climatological base state - iron profiles

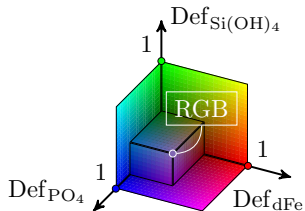
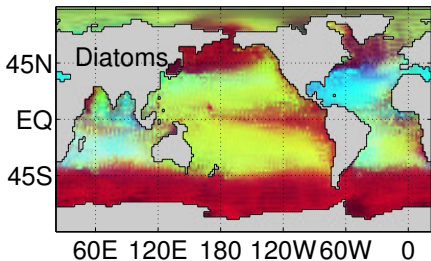
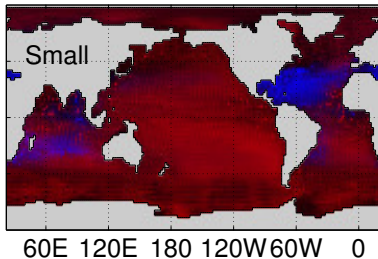
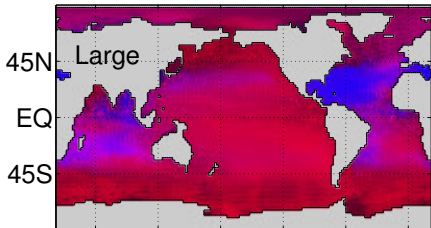
Model with  $\sigma_A = 6 \text{ Gmol}_{\text{dFe}}/\text{yr}$



# Climatological base state - limiting nutrients

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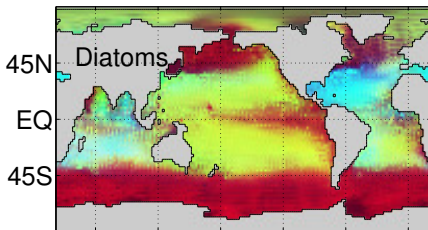
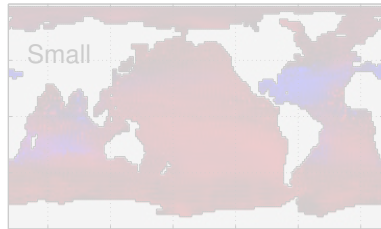
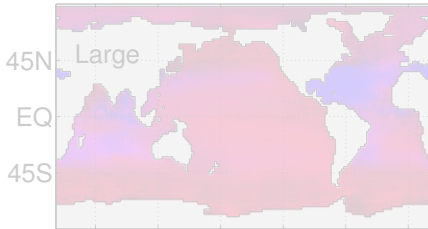
Nutrient Deficiencies for S, L, and D classes



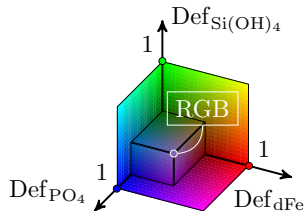
# Climatological base state - limiting nutrients

Model with  $\sigma_A = 6 \text{ Gmol}_{\text{dFe}}/\text{yr}$

Nutrient Deficiencies for S, L, and D classes



60E 120E 180 120W 60W 0

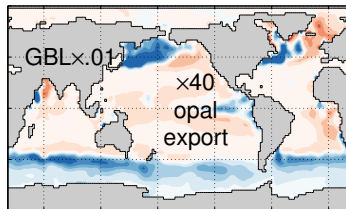
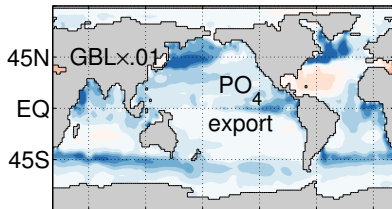


# Export production anomaly due to perturbed aeolian iron

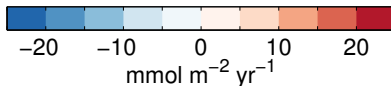
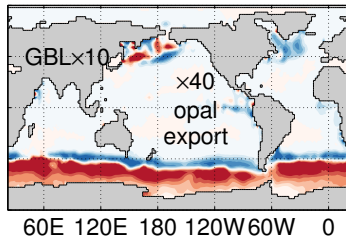
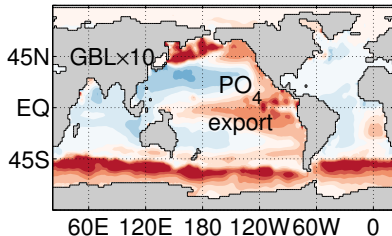
Phosphorus export

Opal export

Globally reduced  
aeolian iron  
( $\times 0.1$ )



Globally increased  
aeolian iron  
( $\times 10$ )



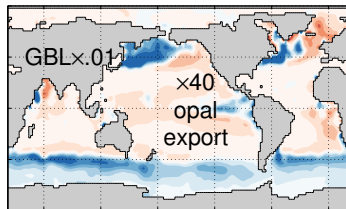
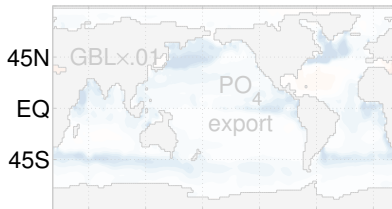


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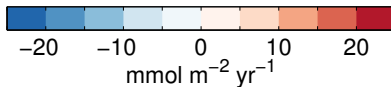
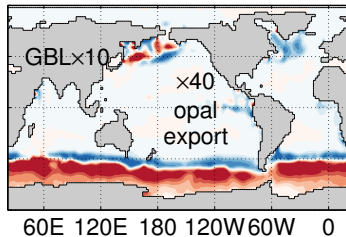
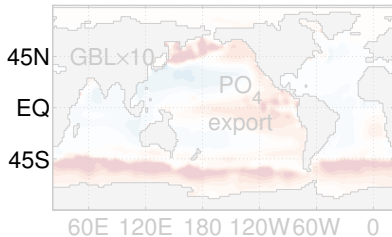
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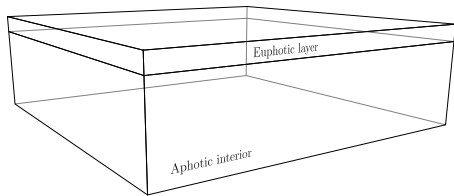
Globally reduced  
aeolian iron  
( $\times 0.1$ )



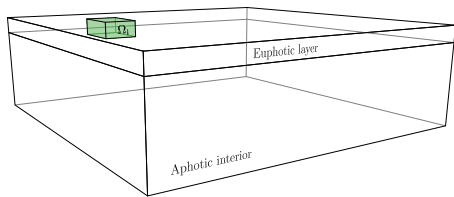
Globally increased  
aeolian iron  
( $\times 10$ )



# Path densities of regenerated nutrients

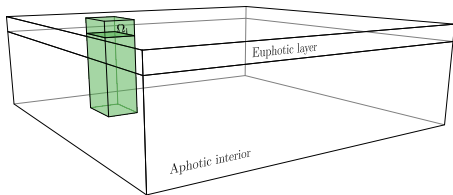


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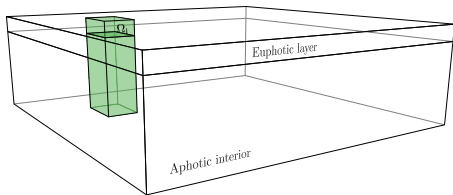
# Path densities of regenerated nutrients

1. Extract nutrient X's regenerated source:  $s_{\text{reg}}^X = \mathbf{S}_{\text{ex}}^X \mathbf{u}^X(\mathbf{x})$



## Path densities of regenerated nutrients

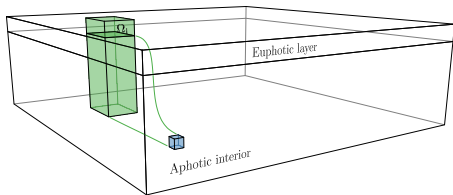
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2. New linear equation:  $(\partial_t + \mathbf{T} + \mathbf{L}_0)\mathbf{x}_{\text{reg}} = \mathbf{s}_{\text{reg}}^X$



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$$(\partial_t + \mathbf{T} + \mathbf{L}_0)\mathbf{g}_{\text{reg}}(t) = \mathbf{0} \quad \text{and} \quad \mathbf{g}_{\text{reg}}(0) = \text{diag}(\mathbf{s}_{\text{reg}}^X)\Omega_i$$



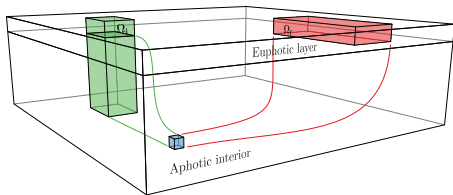
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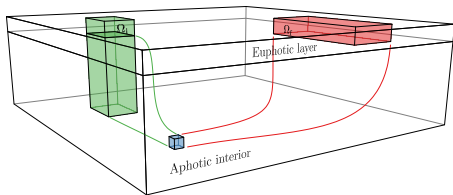
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5. Compute by direct inversion:

$$\langle \mathbf{g}_{\text{reg}} \rangle = (\mathbf{T} + \mathbf{L}_0)^{-1} \text{diag}(\mathbf{s}_{\text{reg}}^X)\Omega_i$$

$$\langle \tilde{\mathbf{G}}_{\text{reg}} \rangle = (\tilde{\mathbf{T}} + \mathbf{L}_0)^{-1} \mathbf{V}\mathbf{L}_0\Omega_f$$





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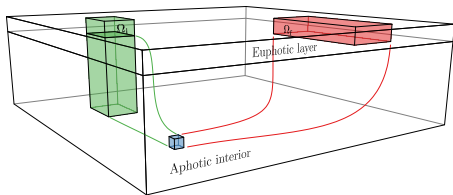
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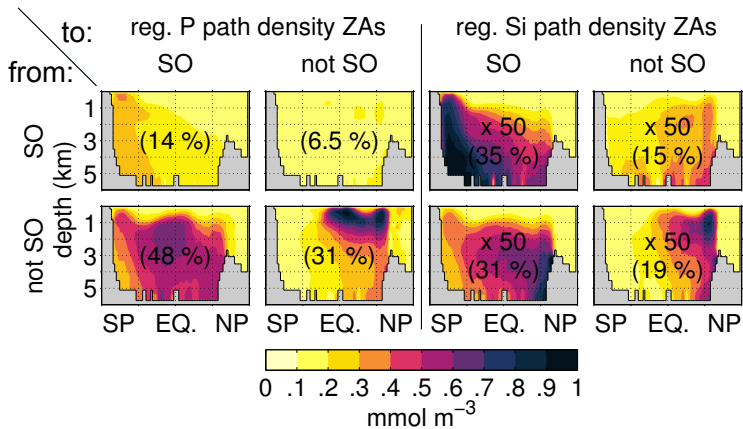
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6. Combine into path density:

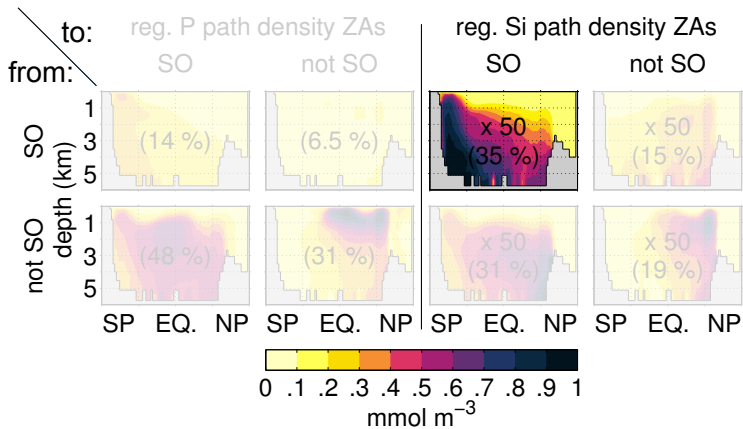
$$\langle \boldsymbol{\eta}_{\text{reg}} \rangle = \langle \tilde{\mathbf{g}}_{\text{reg}} \rangle \odot \langle \mathbf{g}_{\text{reg}} \rangle \quad (\text{element-wise multiplication})$$



# PD of regenerated $\text{PO}_4$ and $\text{Si}(\text{OH})_4$ - base state



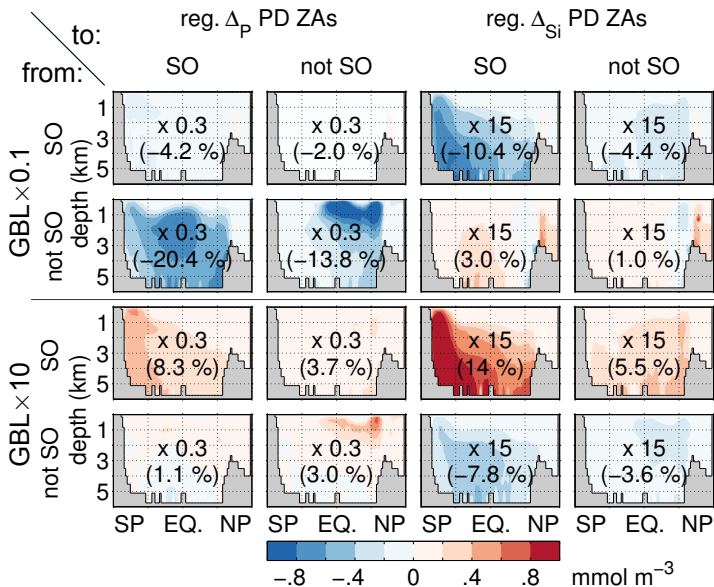
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# PD of regenerated $\text{PO}_4$ and $\text{Si}(\text{OH})_4$ - anomaly

Globally reduced  
aeolian iron

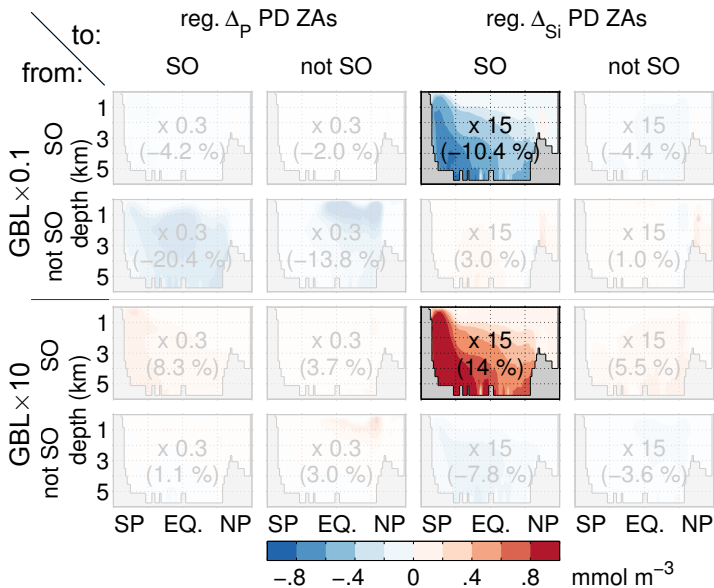
Globally increased  
aeolian iron



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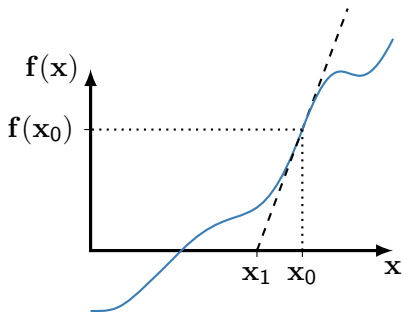
Questions?



## Newton PDE solution

- steady state:  $\partial_t \mathbf{x} = \mathbf{f}(\mathbf{x}) = \mathbf{0}$
- use Newton's Method (generalized zero search)  
linear approximation:

$$\mathbf{f}(\mathbf{x}_1) = \mathbf{f}(\mathbf{x}_0) + \mathbf{Df}(\mathbf{x}_0) (\mathbf{x}_1 - \mathbf{x}_0) + o(\|\mathbf{x}_1 - \mathbf{x}_0\|)$$



where  $\mathbf{Df}$  is the Jacobian,  
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where  $n \sim 600,000!$

To get  $\mathbf{f}(\mathbf{x}_1) \sim \mathbf{0}$ ,  
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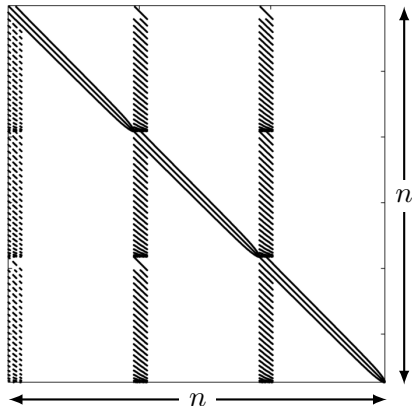
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*Kelley, 2003*

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